

Using initialized simulations to diagnose the growth of systematic biases in the coupled system

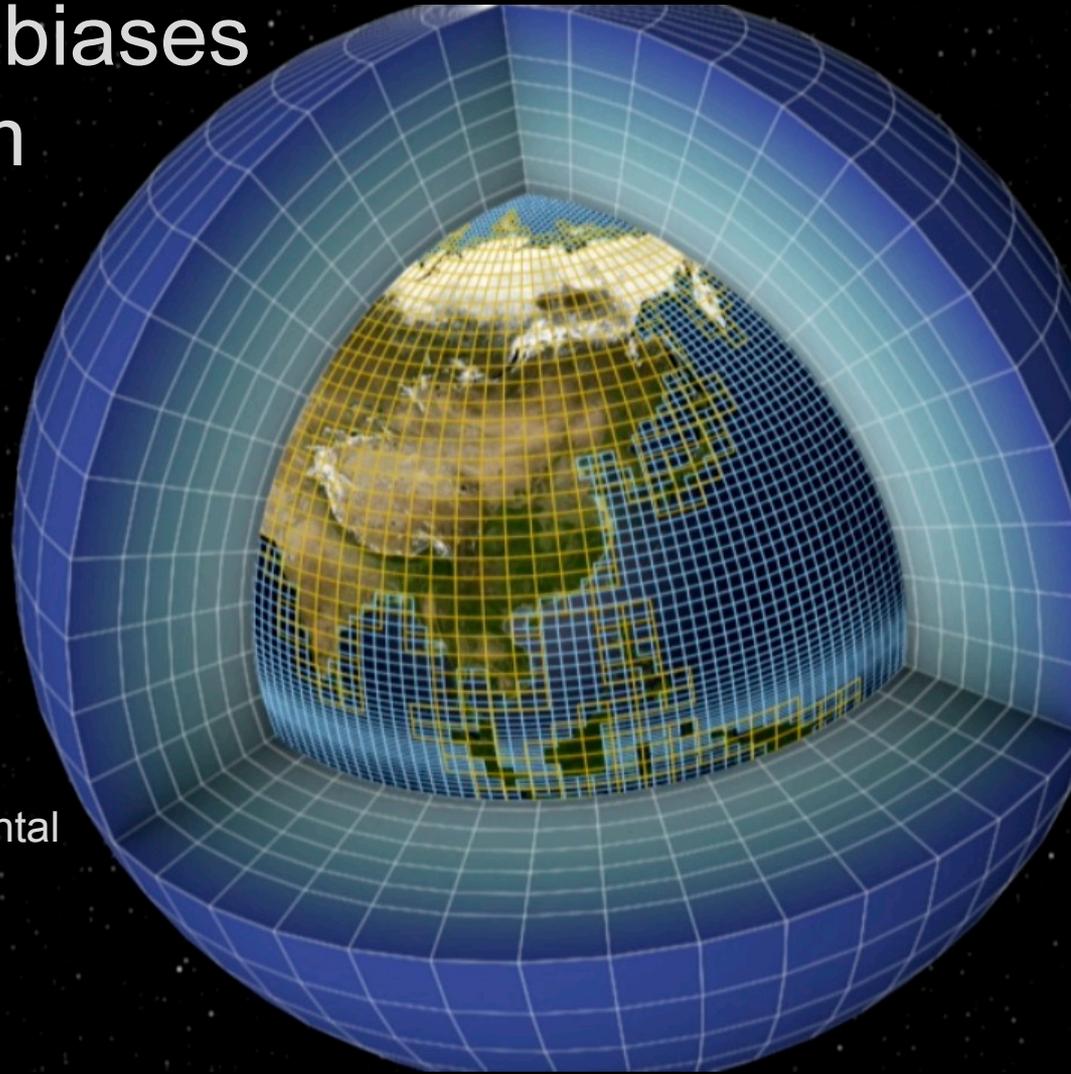
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NCAS-Climate, University of Reading

Outline

- Attributing SST errors to specific components and parameterisations
- Can we devise a systematic experimental approach to guide targeted model development ?
- Discussion points

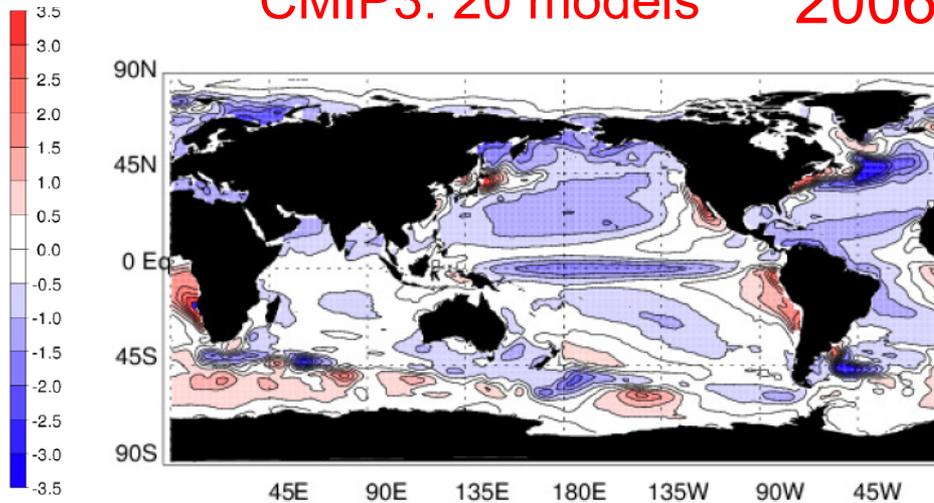


GCMs suffer from systematic and pervasive biases

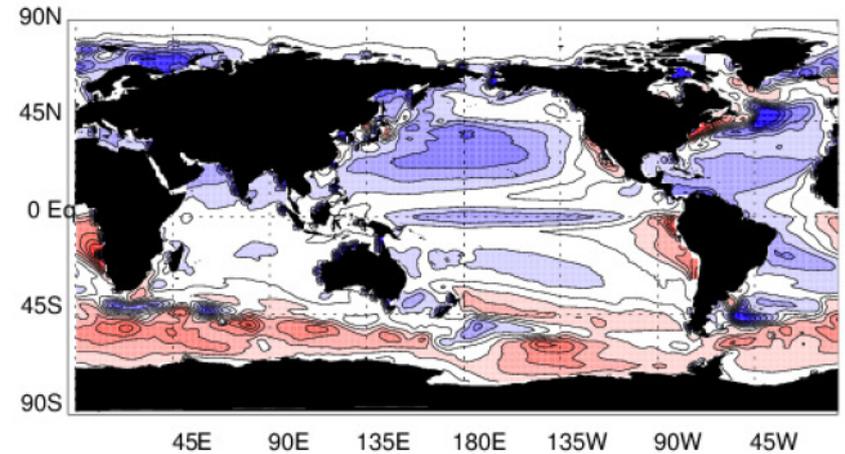
Quiz !

SST errors

CMIP3: 20 models 2006

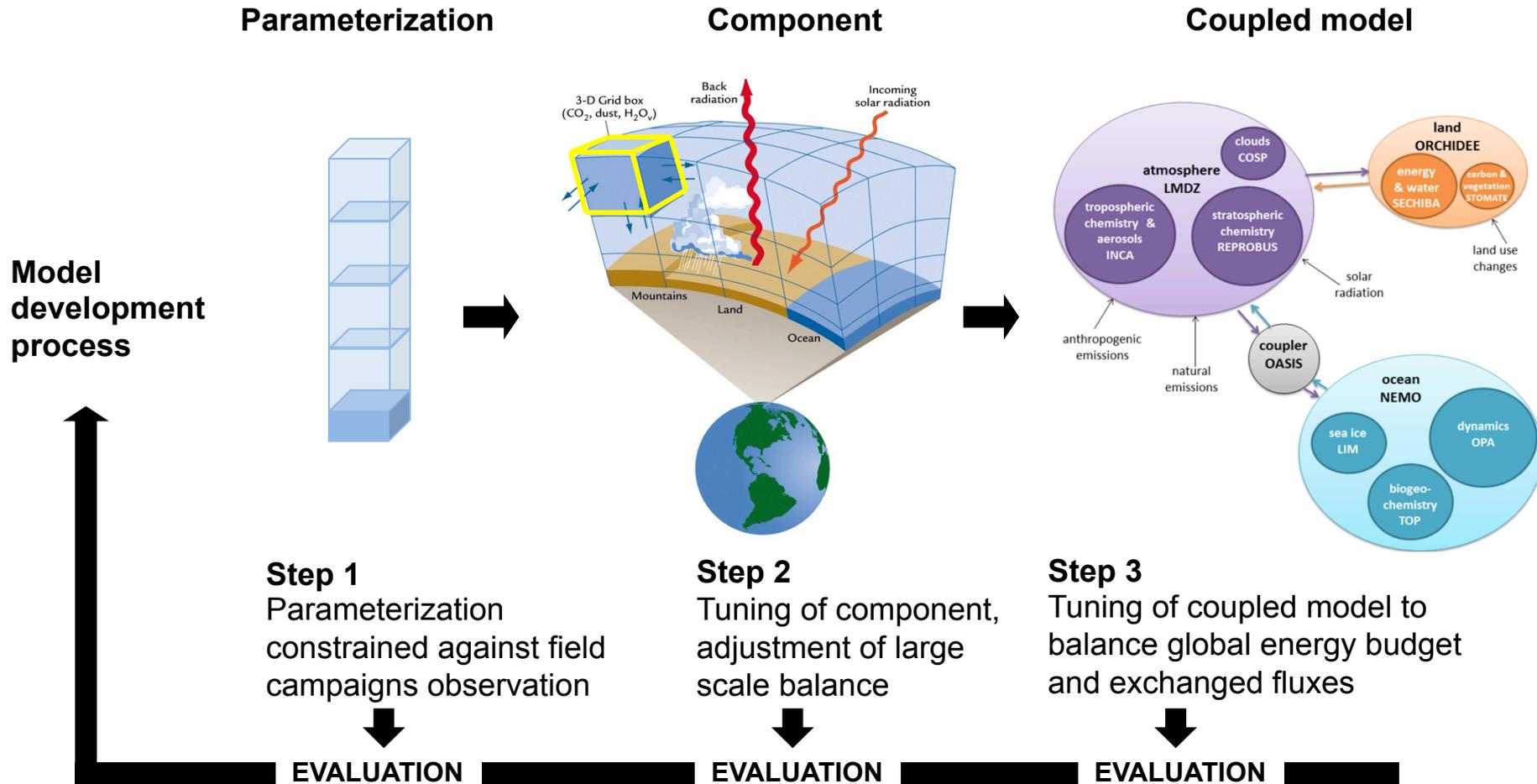


CMIP5: 21 models 2013



- ↪ Severe impacts on climate understanding and prediction
- ↪ WCRP & CLIVAR defined as a priority the understanding of climate models biases
- ↪ **Lack of progress: try out new strategies**

“Classical” CGCMs development path



- ↳ Each step generates biases
- ↳ Source in coupled model is difficult to identify because of **bias compensation**, **feedback amplification** and **non-linearities**
- ↳ This development strategy does not allow to predict the coupled model SST biases

Working backwards

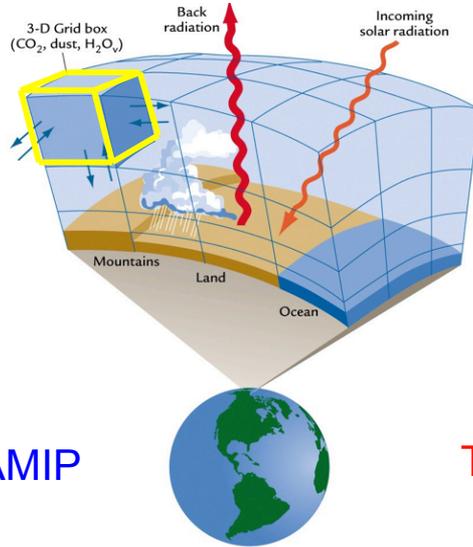
Model development process

Parameterization



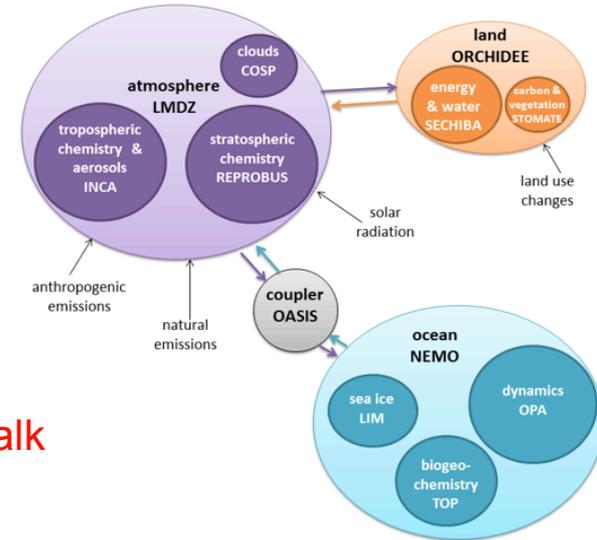
e.g. T-AMIP

Component



This talk

Coupled model



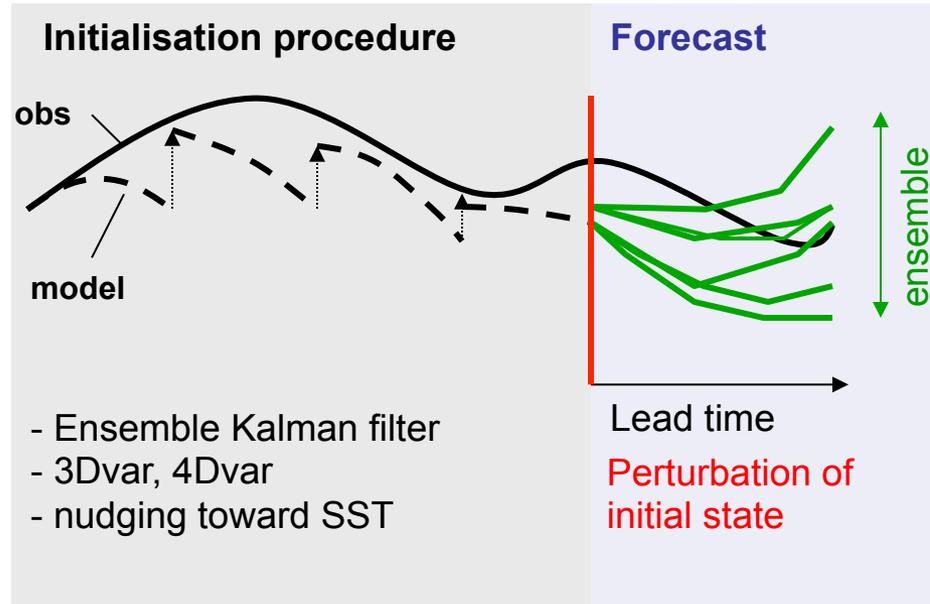
“Fast” physics errors

Atmosphere and ocean component errors

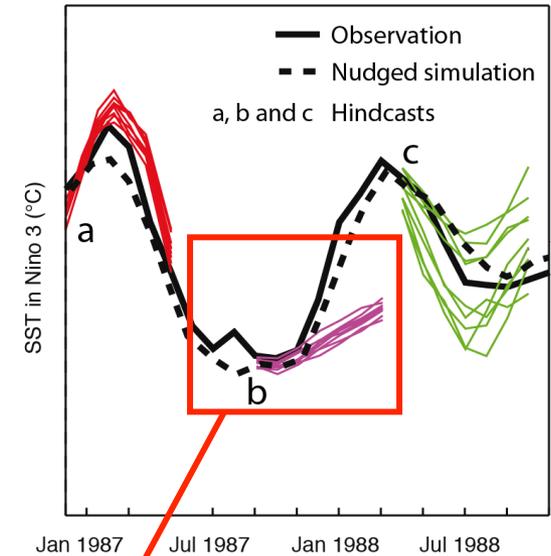
Long term coupled errors

→ Use “reverse engineering” to attribute a **particular** bias of the coupled model to a component and back to a **specific** parameterisation

Using initialised simulations to understand model errors



Hindcasts = forecast of the past period



Adjustment time scale depends on physical processes involved

Hindcasts:

- ↳ Help distinguish time scale and location of error growth
- ↳ Help propose hypothesis for error source

Using coupled hindcast to understand model biases

Previous work on the development of SST errors and attribution to specific component biases:

- Seasonal/decadal time scale:
 - Tropical Atlantic: B. Huang et al. (2007)
 - Tropical Pacific : B. Vanni re et al. (2013, 2014), J. Shonk et al. (2016)
- Decadal/longer time scale:
 - Tropical Atlantic: T. Toniazzo & S. Woolnough (2013)
 - North Atlantic & AMOC : B. Huang et al. (2015)

2 examples

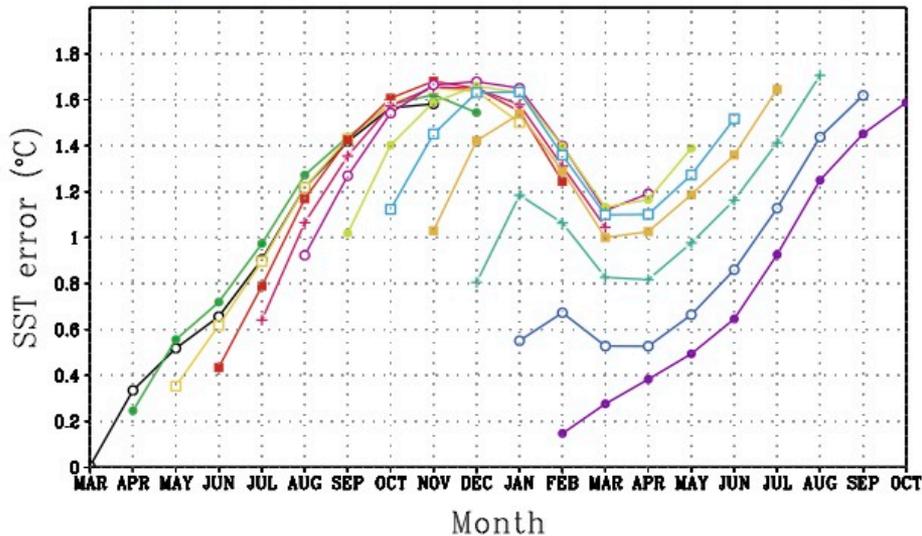
Using coupled hindcast to understand model biases: example 1

Development of SST and D20 errors in the tropical Atlantic

Huang et al. (2007)

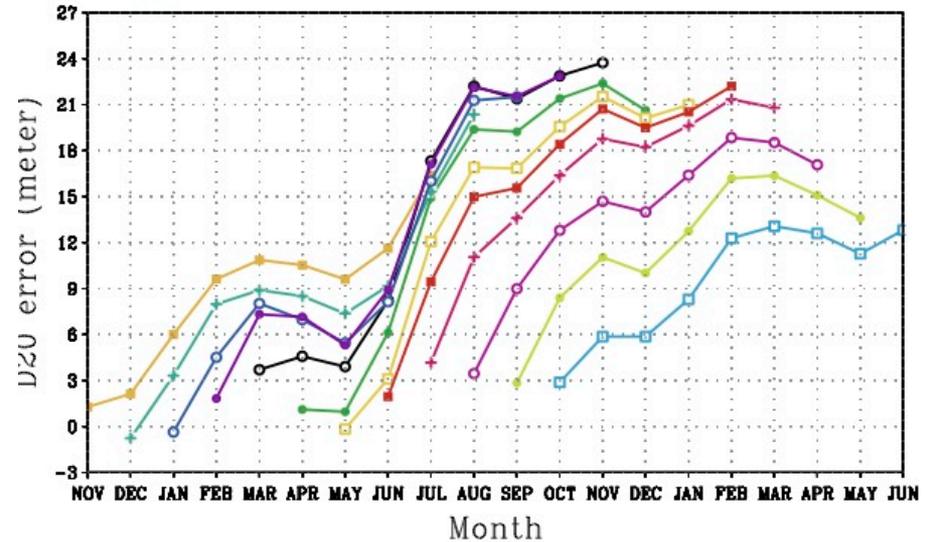
NCEP hindcasts 9 months Lead time

(a) Averaged SST Bias (5°S–15°S, 10°W–10°E)



SST error attributed to heat flux error

(a) Averaged D20 Bias (5°N–10°N, 30°W–50°W)



D20 error attributed to wind stress curl error

✓ Causality is *inferred* as no additional simulations are carried out

Using coupled hindcast to understand model biases: example 2

Development of SST errors in the equatorial Pacific in ECMWF system 4



- Cooling bias in west Pacific due to too strong trade winds
- Error appears in first 10 -20 days
- Exact same behaviour in uncoupled mode
- Model developers can focus analysis on uncoupled

Using coupled hindcast to understand model biases

Previous work on the development of SST errors and attribution to specific model biases

- Seasonal time scale:
 - Tropical Atlantic: Huang et al. 2007
 - Tropical Pacific : Vanni re et al. 2013, 2014, Shonk et al. 2016
- Decadal time scale:
 - Tropical Atlantic: Toniazzo & Woolnough 2013
 - North Atlantic AMOC, role of fresh water flux : Huang et al. 2015

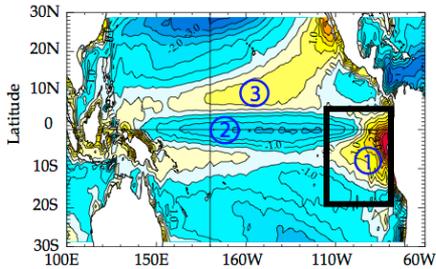
Mechanisms that initiate SST drift usually associated to atmosphere

- Heat fluxes, wind stress (zonal, meridional, curl), fresh water flux

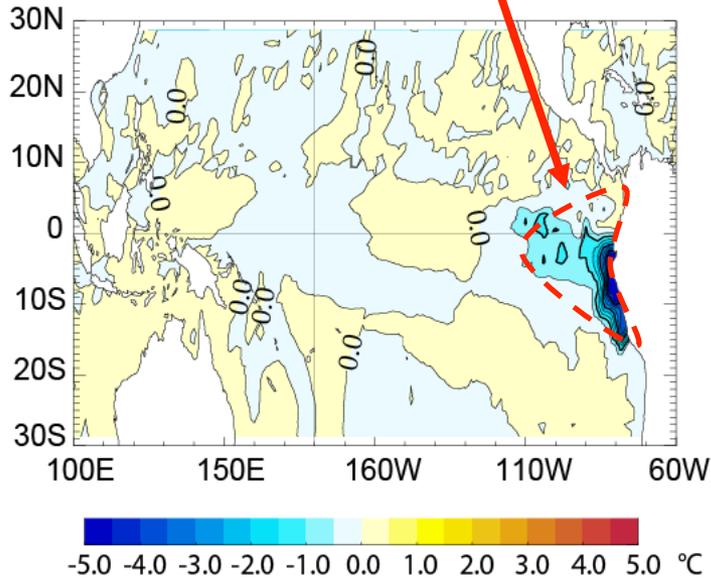
Two types of studies:

- *Infer* causality via diagnostics and physical understanding
- Actually *show* causality via additional simulations

Using additional simulations to demonstrate the source of error

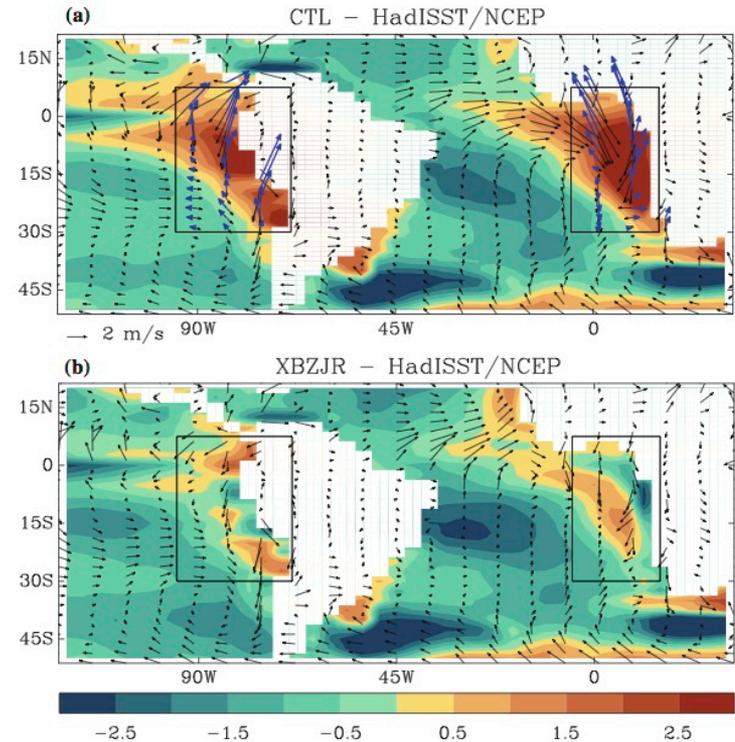


Oceanic simulation forced with fixed flux : impact of meridional wind correction



Vanni re et al. (2014)

Coupled simulation with wind correction



Toniazzo and Woolnough (2013)

Can we devise a systematic experimental approach ?

Vannière et al (2014) proposed a systematic approach to investigate the root cause of a SST bias in a climate model

5 steps for 'solving the case':

1. Identify the **location and seasonality** of the SST bias
2. Examine the **time scales** over which errors develop in different variables and link them together to build a **chain of causality**
3. Find whether the origin of the bias is **local or remote**
4. Determine if an **atmospheric field or an oceanic field** is at fault
5. Investigate whether the error is caused by the **direct effect** of that field, or by **coupled feedbacks**

Associated experiments in support of approach

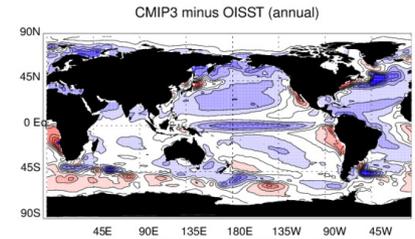
Vanni re et al. (2014)

The 5 steps

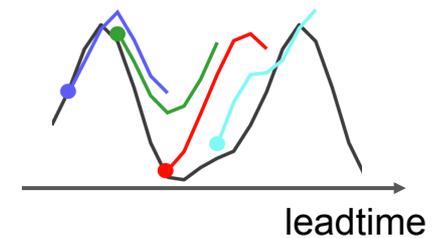
- S1** Location / seasonality
- S2** Time scale / chain of causality
- S3** Local or remote
- S4** Atmospheric / oceanic field responsible for the bias
- S5** Direct effect / amplification by coupled feedbacks

Associated experiments

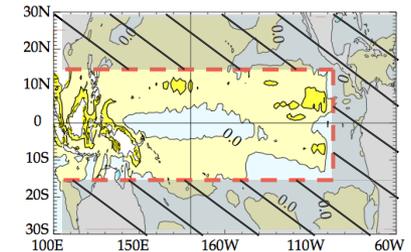
Historical or control experiment



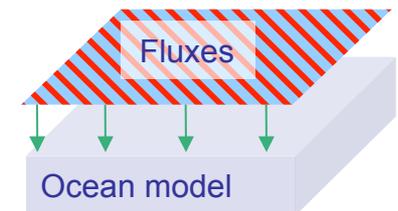
Seasonal to decadal hindcasts



Regionally restored experiments



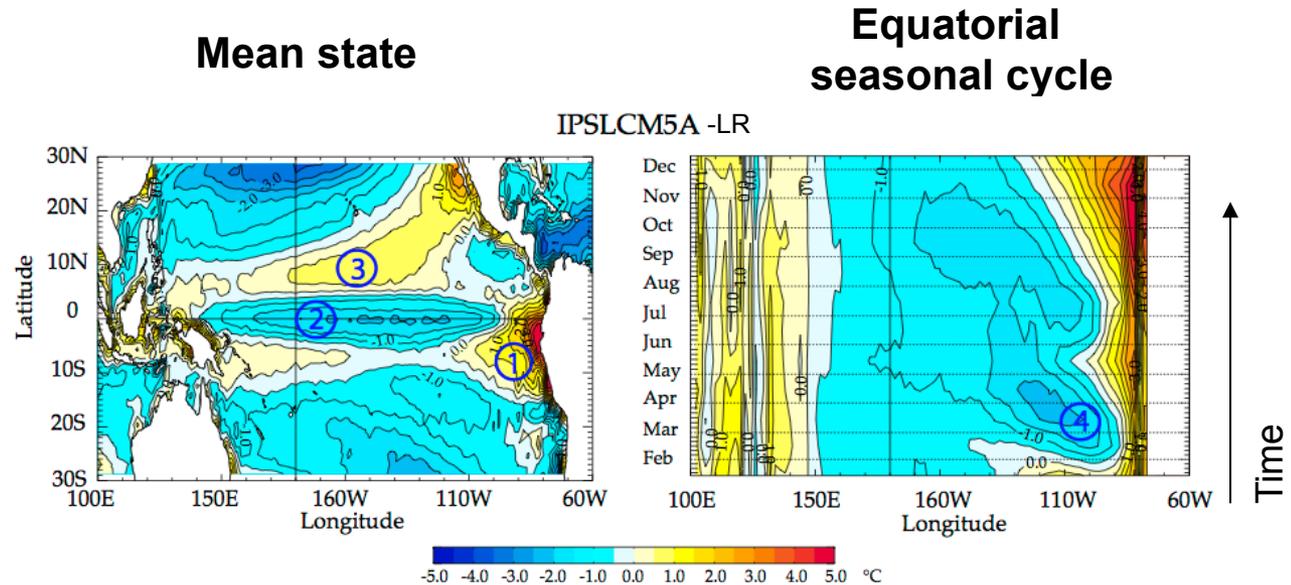
Ocean-only forced experiments



Identifying the origin of SST mean state biases in the tropical Pacific in IPSLCM5A-LR

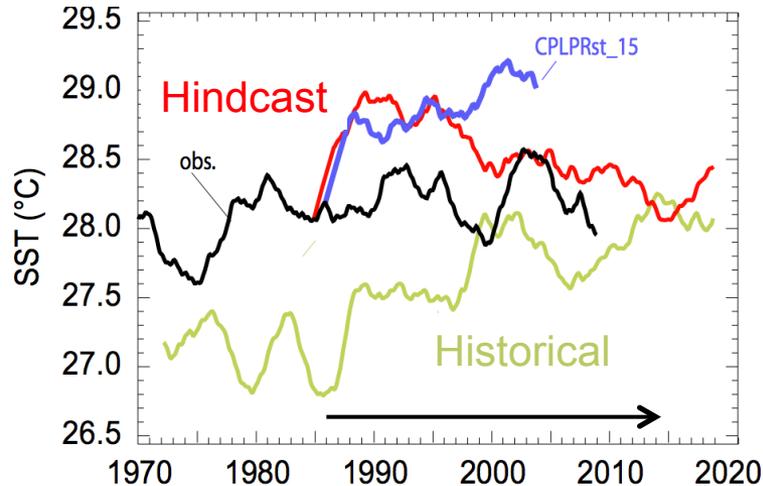
Approach is applied to **cold tongue bias** in IPSL-CM5A-LR (S1)

- ① Warm bias in the east Pacific
- ② **Cold tongue bias**
- ③ Warm bias on both side of the equator
- ④ Spurious spring upwelling bias

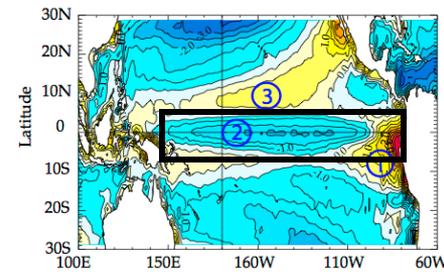


Cold tongue bias origin

> **S2** : Time scale → Cold tongue bias



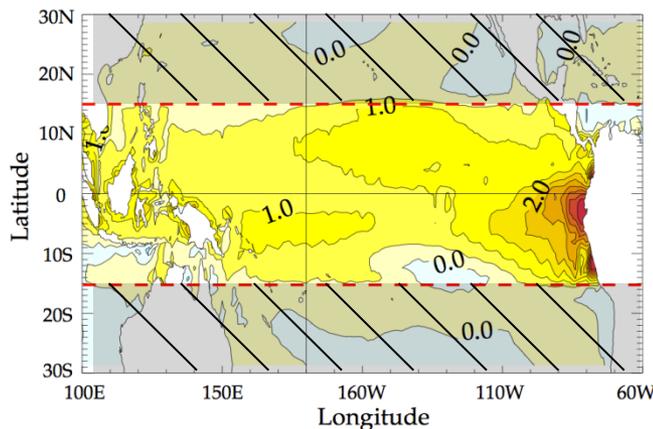
- ✓ It takes 30 years for the cold tongue bias to appear at the equator
- ✓ Hypothesis : ocean slow dynamics



> **S3** : Geographical origin → Cold tongue bias

CPLPrst_15: Initialised simulation restored toward observed SST in midlatitudes

 SST nudging

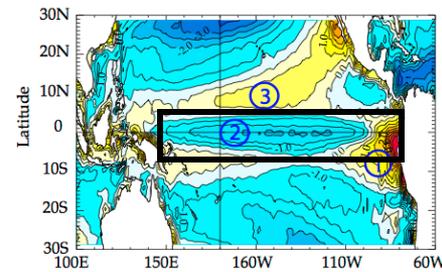


20-yr leadtime

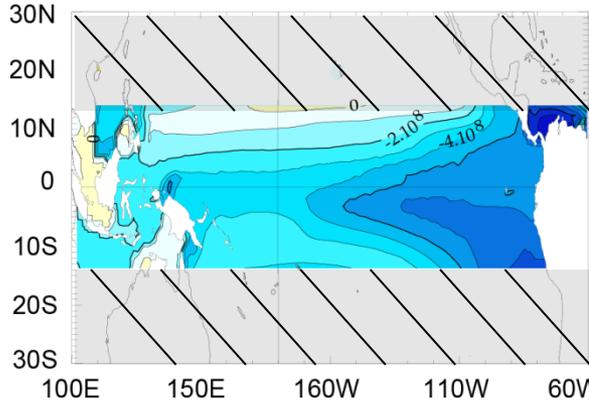
- ✓ SST corrected in mid-latitudes
- ☒ no development of the cold tongue bias

Cold tongue bias origin

> **S4** : Ocean only simulation → reproduce cold tongue bias



Ocean only simulation (bulk) 300m heat content



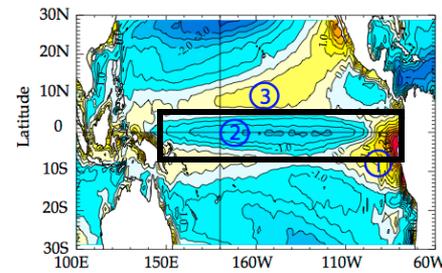
Nudging toward the hindcast mid-latitudes cold bias

30 years leadtime

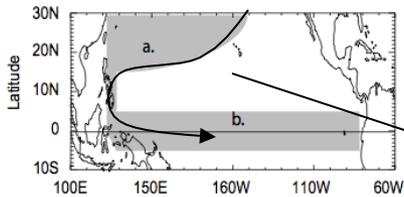
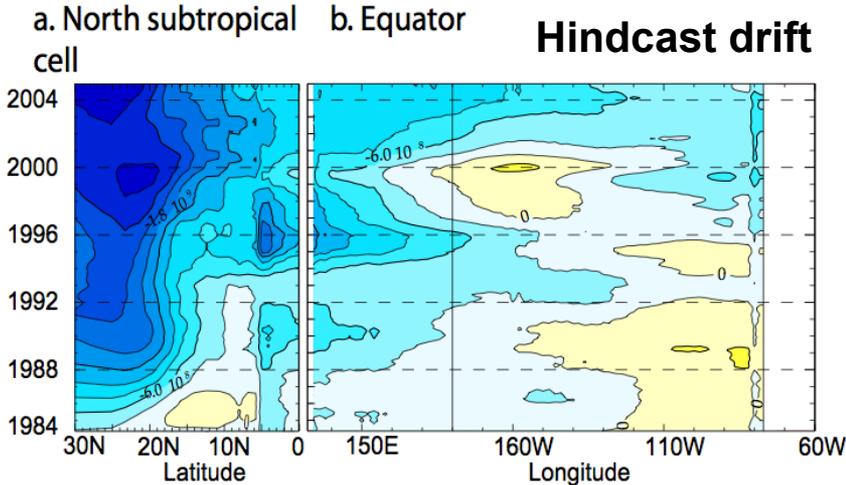
Experiment	Equatorial cooling trends of the 300m HTC ($\text{J} \cdot \text{m}^{-2} \cdot \text{mth}^{-1}$)
Hindcast	$-1.35 \cdot 10^6$
Ocean only with SST restoring	$-1.73 \cdot 10^6$
Ocean only without SST restoring	$0.1 \cdot 10^6$

- ✓ When the midlatitudes cold SST bias is prescribed in an ocean-only experiment, the cold tongue bias develops at the equator
- ✓ The cooling trend is similar to that simulated by the control hindcast

Cold tongue bias origin



A possible cause of the midlatitude cold bias propagation is the advection by subtropical cells



Subtropical cells path according to Izumo et al. (2002)

Differs from other sources of the cold tongue bias (Vannière et al. 2013)

- Bjerknes feedback (Met Office)
- Atmospheric component wind errors (INGV)
- or otherwise proposed in many studies

Summary and discussion points (1/2)

- New approaches needed to address SST systematic errors
- Strategy to relate **coupled errors** to the **errors in one component** independently of the coupling:
 - 5 step 'case solving' approach
 - Requires range of dedicated simulations, including initialized
 - Proof of concept from several studies (tropical Pacific and Atl.)
 - Further benefits/costs to explore:
 - Apply during model development phase
 - cheap (300 years)
 - but need to develop a 'tool box', i.e. several types of simulations (one time investment)
 - Precise types of simulations will depend on 'case' i.e. SST bias – no 'standard' set
 - Can't be directly applied to SST interannual variability biases (ex: ENSO) but can be applied to ENSO mechanisms and feedbacks (not shown)

Summary and discussion points (2/2)

- But, AMIP/T-AMIP is the starting point in the tropics
 - Most SST errors initially due to fast atmosphere biases
- Decadal bias investigation : start with simplified forcing (better signal/noise ratio)
- Prospects for high-resolution to ameliorate climate biases
 - No systematic impact of HR, although dynamics better
 - Works together with physical parameterisations
 - Coherence of air-sea interactions and horiz./vert. resolution
- Atmosphere and ocean biases that lead to SST, SSS bias
 - Fluxes, usually an atmosphere cause
 - Near the equator: wind stress, wind stress and wind stress
 - Data assimilation won't help
 - Mid-latitudes: subduction in the ocean, fresh water fluxes

Ocean-only forced experiments

Bulk formulation

Computed interactively with low level atmospheric fields

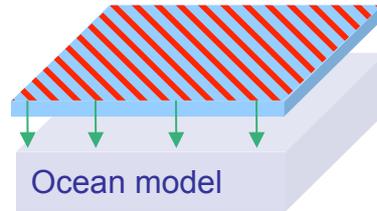
Coupled model

Ex: 10M wind

Observation

Ex: (t2m, q2m, SW, LW..)

DFS4.3, Brodeau et al. 2010



- damping of SST errors
- dynamical and thermodynamical effects of wind

Fixed flux formulation

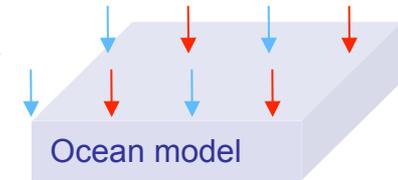
Air-sea fluxes prescribed

Coupled model

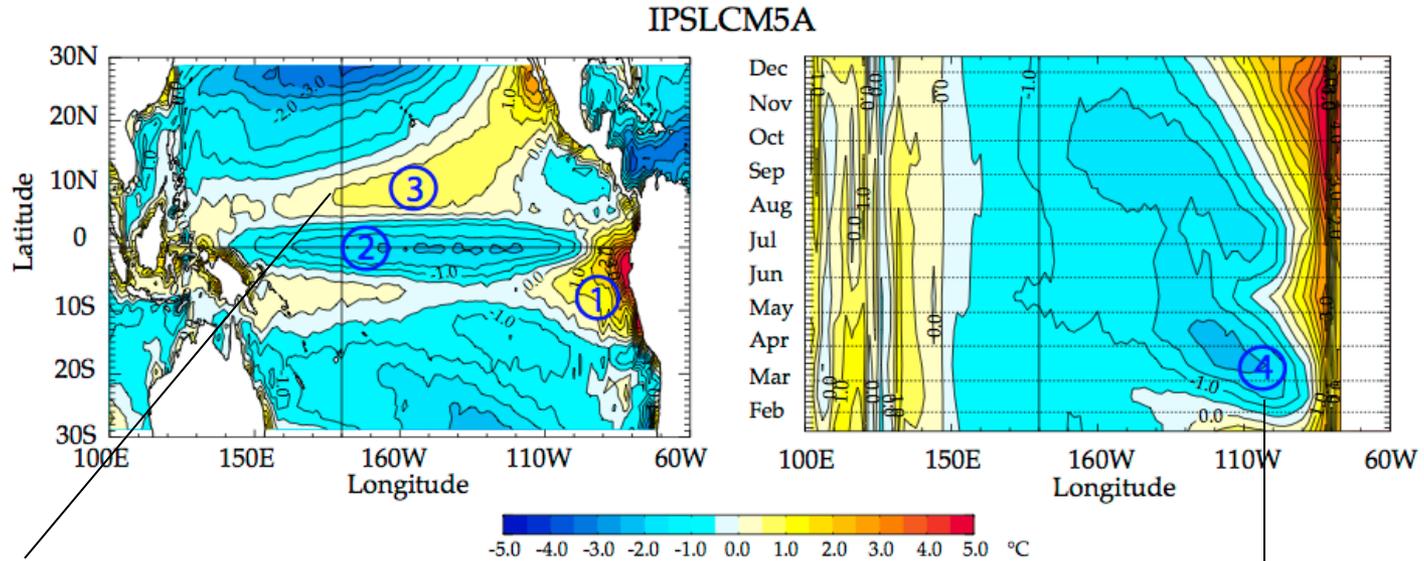
Ex: Wind stress

Observation

Ex: (solar HF, non-solar HF...)



- development of SST errors
- dynamical effect only of wind



③

Hindcast : development in ~ 6 months

Ocean-only simulations :

- too strong SW heat flux
- too low latent heat flux
- no dynamical role of the wind

④

Hindcast : development in ~ 6 months

Ocean-only simulations :

- coupled bias
- EP warm bias modifies the wind seasonal cycle and initiate the upwelling

Cold tongue bias :

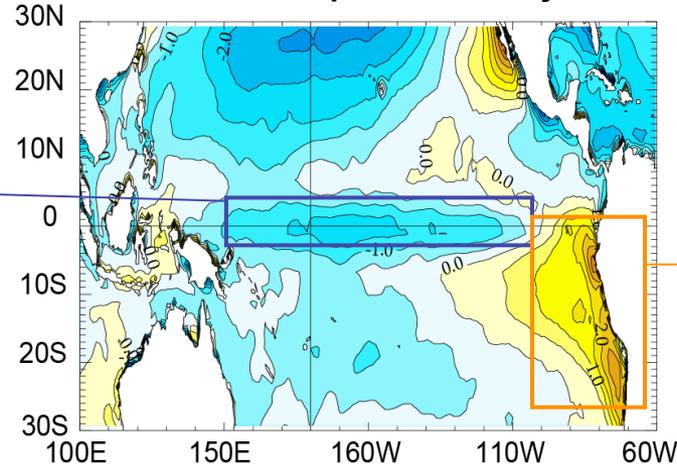
- too strong easterlies at the equator

- amplification by coupled feedbacks such as Berknes feedback (Dijkstra and Neelin 1995, Lin 2007)

- too diffuse thermocline (Davey et al. 2002)

...

CMIP5 SST compared with Reynolds



East Pacific warm bias :

- meridional and coastal wind

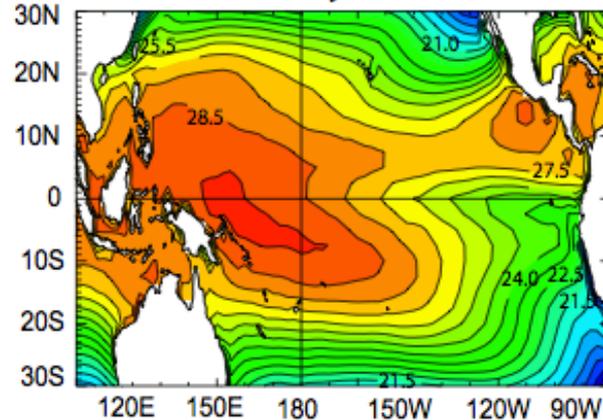
(deSzoek et Xie 2008, Large et Danabasoglu 2006)

- low level clouds (Mechoso et al. 1995, Ma et al. 1996, Wang et al. 2004)

- eddies (Colas et al. 2012)

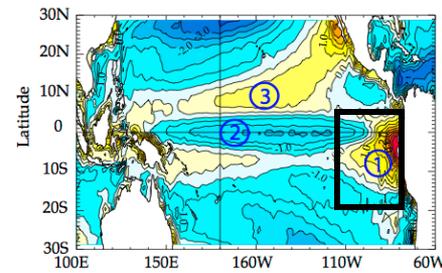
- possible interaction with the double ITCZ bias

Reference Reynolds SST

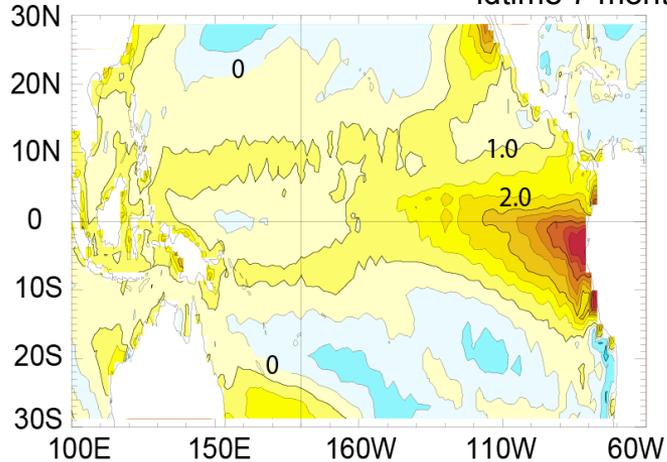


Origin of the East Pacific warm bias

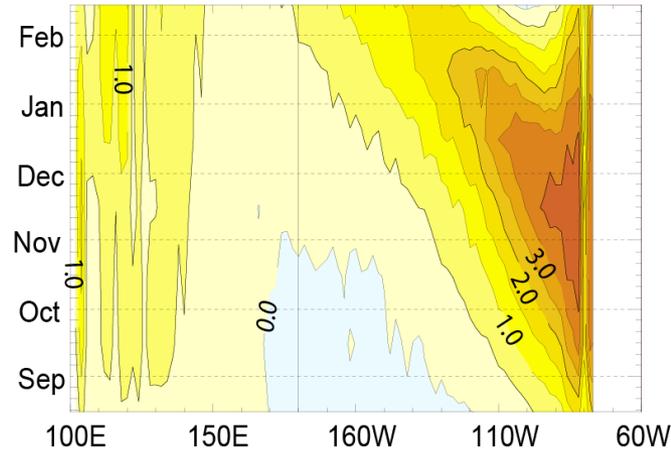
> S2 : Time scale and propagation → EP warm bias



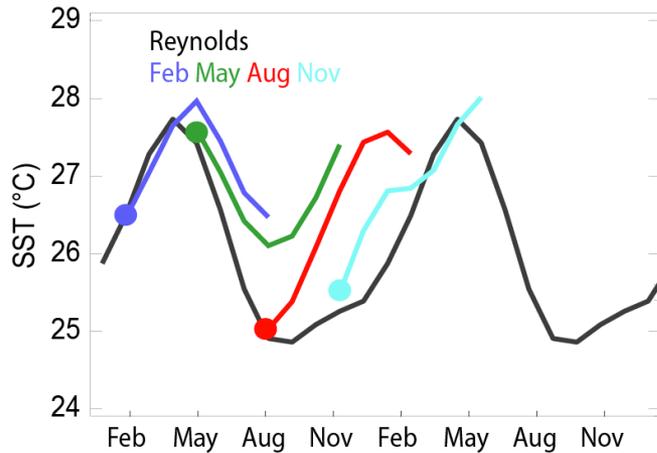
a. Mean state SST drift Init in Aug. ldttime 7 month



b. Equatorial SST drift Init in Aug. ldttime 7 month



c. Nino 3 SST

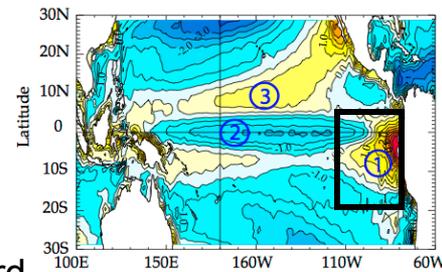


✓ Development of the warm bias at the Peruvian coast and propagation toward the west

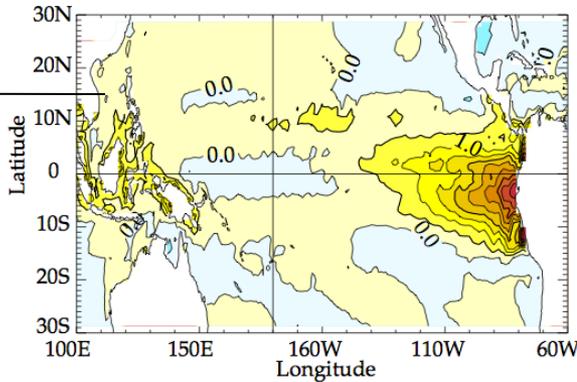
✓ Development of the warm bias whatever the start date is and maximum development of the bias during the upwelling period.

Origin of the East Pacific warm bias

> S3 : Geographical origin → EP warm bias

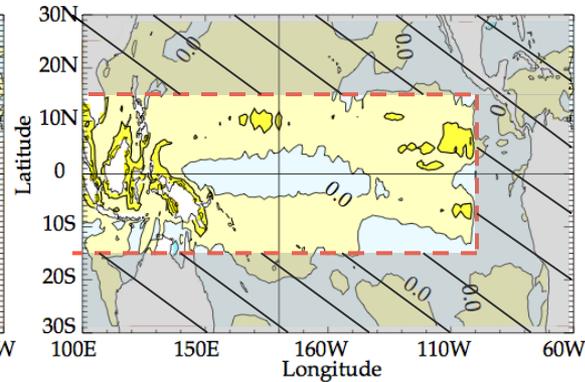


Hindcast + wind from the simulation nudged toward SST

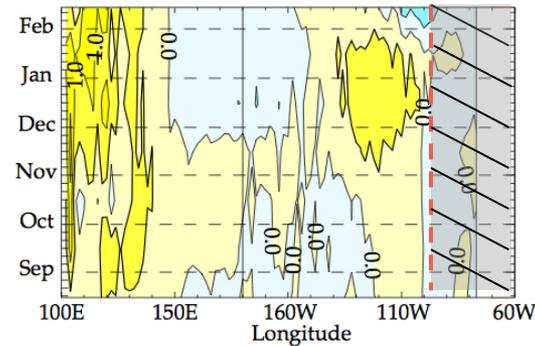
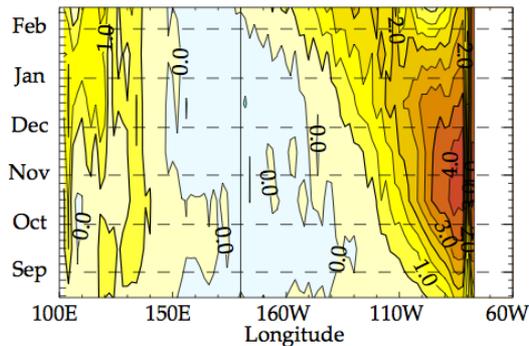


No SST bias feedback on the surface wind

Hindcast + nudging toward SST



 SST nudging



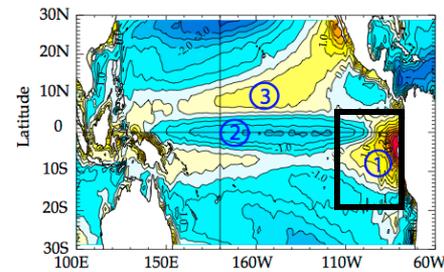
The amplification of the warm bias doesn't require a dynamical coupling

Advection is a key process for the westward propagation of the bias

✓ The warm bias is generated at the Peruvian coast and is advected by oceanic currents toward the west.

Origin of the East Pacific warm bias

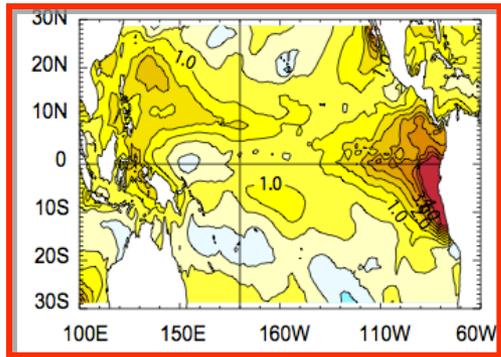
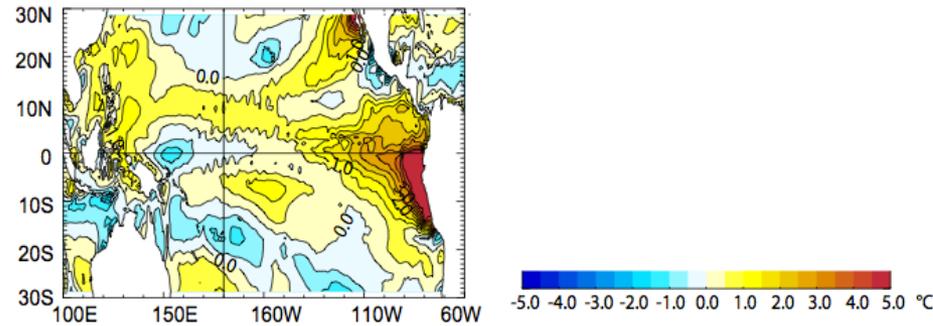
> S4 : Field responsible for the bias → EP warm bias



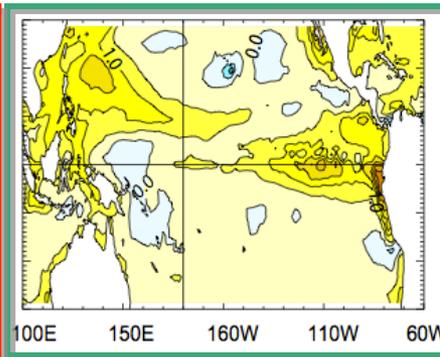
SST errors in hindcasts
3 month ldtme



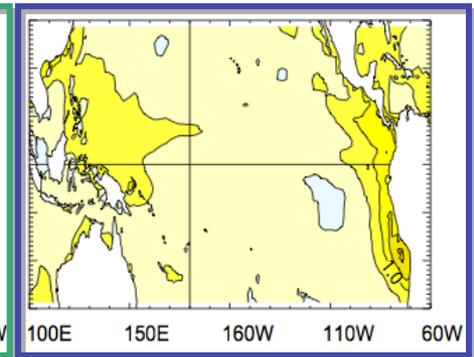
Ocean-only simulations
(bulk)



All field from the hindcast



10m wind impact



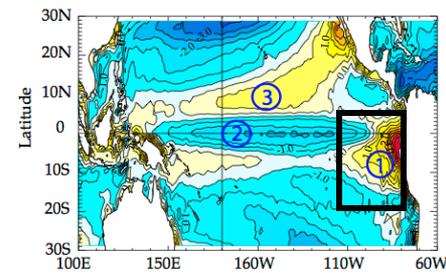
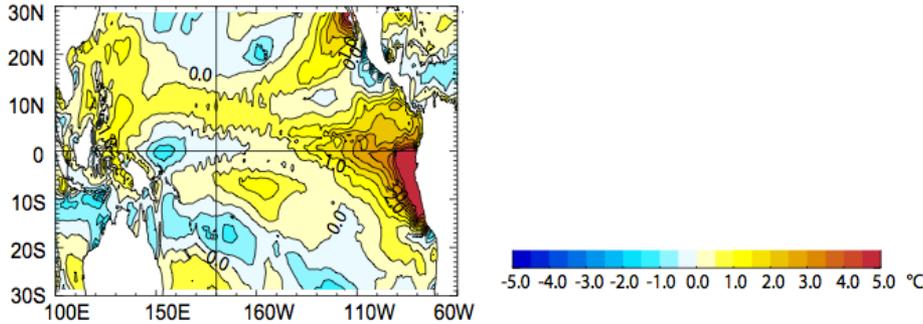
SW impact

- ✓ Ocean-only simulation forced by hindcasts fields decomposition of the different contributions to SST biases
- ✓ 10m wind represent the main contribution to the warm bias

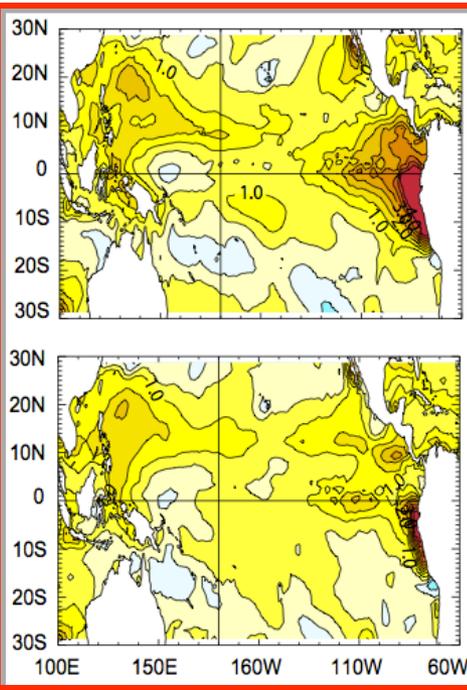
Origin of the East Pacific warm bias

> S5 : Degree of coupling → EP warm bias

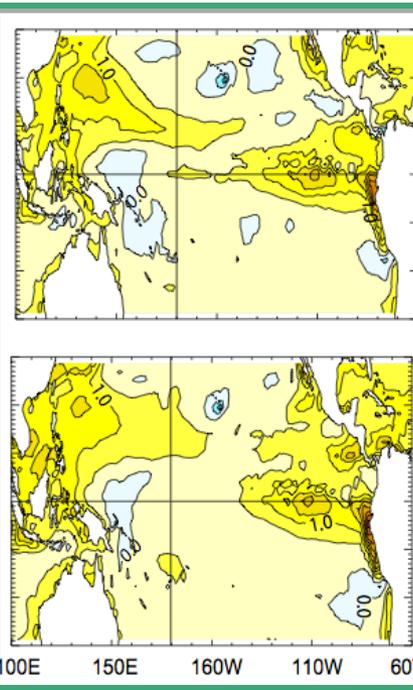
SST errors in hindcasts after a 3 month lead time



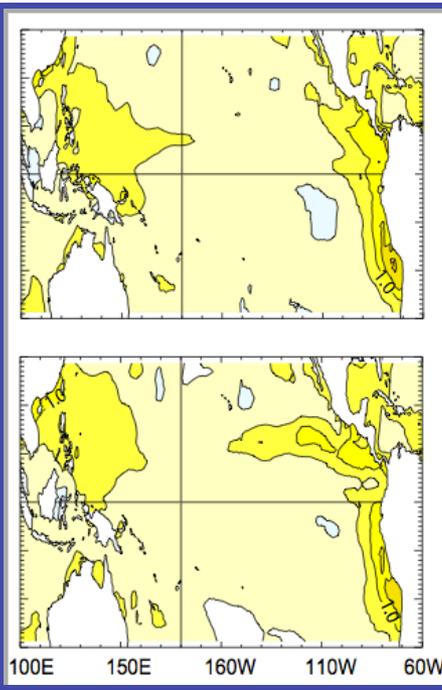
Hindcast forcing



All field coming from the model



10m wind impact

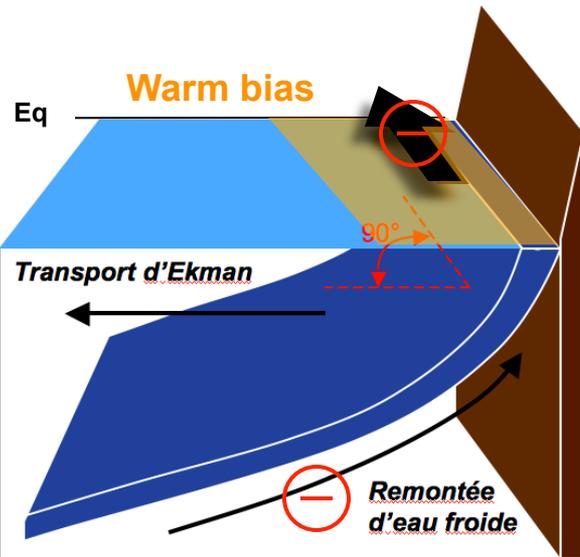
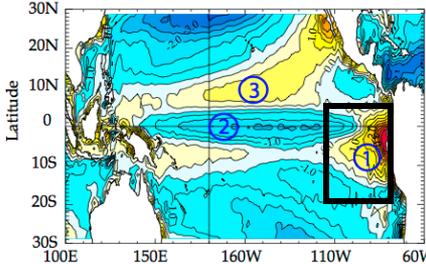


SW impact

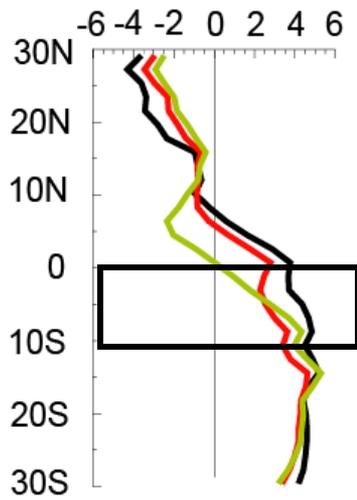
AMIP forcing

✓ EP warm bias is due to a direct effect of the atmosphere component error and the coupling has little impact.

Origin of the East Pacific warm bias

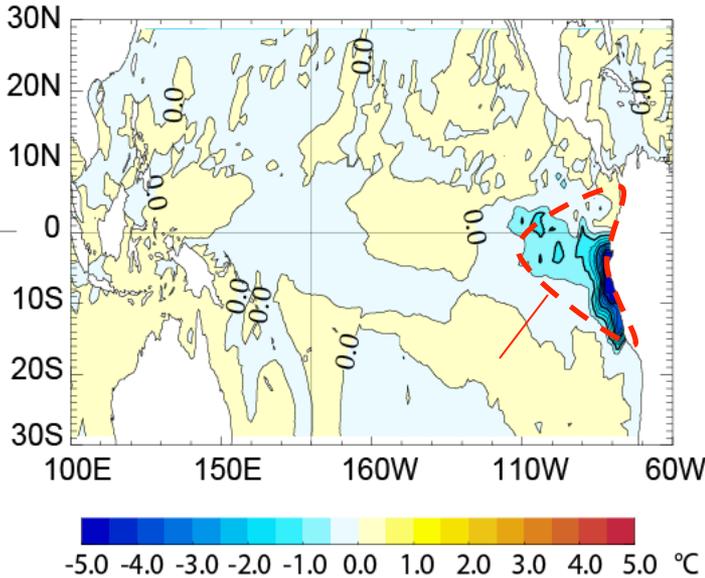


Meridional wind along the Peruvian coast



Observation
AMIP
Historical

Oceanic simulation forced with fixed flux : impact of meridional wind correction



✓ East Pacific warm bias = meridional wind component at the Peruvian coast

Meridional wind is corrected with DFS4.3