

# Tropical Pacific Observing System (TPOS) Equatorial Pacific Experiment (TEPEX)

Contributors: NOAA CPO/CVP, GOMO, OAR (PMEL, PSL, GFDL)  
and NCAR, CU, CSU, SIO, UW, NASA

## Key Points

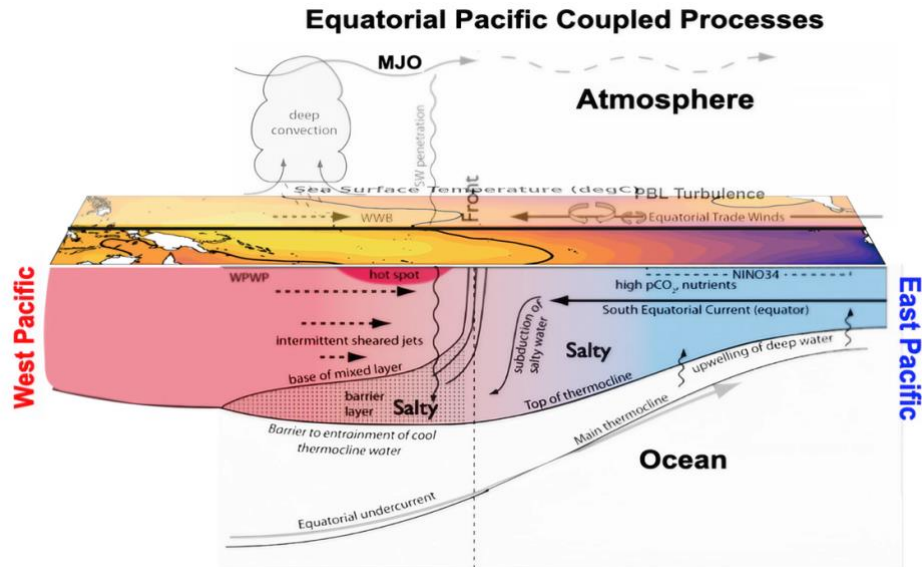
- The tropical Pacific plays a central role in global weather and climate. ENSO, the strongest interannual climate signal, and the MJO, the strongest intraseasonal signal, both shape high-impact environmental events worldwide. ENSO and the MJO interact with each other in the tropical Pacific. Accurate detection, description and modeling of both phenomena is critical to subseasonal to interannual predictions.
- Among the important processes of ENSO, two stand out as inadequately understood and poorly represented by numerical models. One is equatorial upwelling and mixing in the eastern Pacific cold tongue, the other is the zonal movement of the western Pacific warm pool. Both involve active air-sea interaction. They are the primary targets of TEPEX.
- TEPEX emerged from wide discussions in the science community and was recommended by TPOS 2020 as an urgently needed action to advance our understanding and prediction of global variability from subseasonal to interannual timescales, and to inform the evolution of the sustained observing system. TEPEX provides opportunities for multidisciplinary studies of the tropical ocean and atmosphere. It is envisioned as an internationally collaborated and coordinated program and serves as the first field campaign of the WCRP Global Precipitation Experiment (GPEX).
- TEPEX is planned for 2026 - 2028. Its preliminary plan capitalizes on major ongoing upgrades to the TAO moored array and new observing technologies, and has been focused by several years of pre-field modeling studies.
- The field observations of TEPEX in regions without previous comprehensive air-sea interaction field campaigns will enable the global research and operation communities to tackle physical processes key to ENSO prediction through improvement of fundamental understanding and prediction models.

## Motivation

ENSO and the MJO are strong influences on global environmental extremes including precipitation, wildfire, tropical cyclones, flooding, droughts, and heat waves. They also influence marine biological productivity, air quality, and the global carbon cycle. Accurate predictions of ENSO and the MJO are essential to mitigate the societal impacts of multiple environmental events in a changing climate.

Our current prediction models, however, do not adequately represent processes key to ENSO and the MJO, and their interactions. Prediction models misrepresent the cold tongue and warm pool in ways that have far-reaching effects on global precipitation, energy and carbon budgets, and circulation patterns. In consequence, ENSO and MJO prediction skills are limited and do not meet societal needs. Improvement of ENSO and MJO prediction must be based on a solid understanding of their central mechanisms, to enable implementation of these factors into the global models. Detailed in situ observations from targeted field campaigns are the foundations for such understanding.

47 Variability of the East Pacific cold tongue and West Pacific warm pool dominate the zonal  
 48 sea surface temperature (SST) gradient that plays a central role in ENSO and helps determine  
 49 the eastward extent of MJO events. Accurate representations of processes controlling the  
 50 variability of the cold tongue and the warm pool are critical for enhancing prediction skills at the  
 51 global scale and intraseasonal to interannual timescales. Both the maintenance and variability of  
 52 the cold tongue and warm pool are actively involved with air-sea interaction (Fig. 1) through the  
 53 air-sea transition zone, which includes the upper ocean, air-sea interface, Marine Atmospheric  
 54 Boundary Layer (MABL) as a single identity. Rainfall, surface winds, and modulation of solar  
 55 radiation by clouds govern the input of freshwater, momentum, and energy to the ocean. The  
 56 oceanic response to these disturbances is regulated through the upper-ocean stratification of  
 57 temperature, salinity, and velocity which determine the extent to which the surface inputs  
 58 penetrate vertically and are transported horizontally. The consequent distributions of the upper-  
 59 ocean heat content and sea surface temperature (SST) feed back to the atmospheric wind,  
 60 clouds, and rainfall. Strong horizontal gradients in SST exist in both the cold tongue (in the  
 61 meridional direction) and the eastern edge of the warm pool (in the zonal direction). The air-sea  
 62 interaction processes uniquely acting upon the strong SST gradients in these two regions cannot  
 63 be fully understood based on data obtained from elsewhere.



64 Figure 1. Illustration of air-sea interaction processes of the equatorial Pacific (From Brown et al. 2014)

65

66 **Goal and Objectives**

67 The overarching goal of TEPEX is to enhance our understanding of the key processes for  
 68 the intraseasonal to interannual variability of the tropical Pacific, especially those that govern the  
 69 evolution of ENSO, thereby to provide robust information guiding improvement of ENSO  
 70 prediction. This goal will be achieved through three related efforts: (1) Conduct field observations  
 71 targeting two regions: the equatorial cold tongue of the eastern Pacific and the eastern edge of  
 72 the warm pool of the central Pacific, (2) Combine observational analysis and numerical modeling  
 73 to dissect the detailed processes critical to air-sea interaction of the equatorial Pacific and ENSO  
 74 dynamics, and (3) Apply these gains to model advancement: new parameterizations of ocean  
 75 vertical mixing, and explicitly-modeled freshwater- salinity coupling.

76

77 **Physical basis**

78 The physical processes of the cold tongue and warm pool share many commonalities  
 79 (e.g., strong surface winds, air-sea fluxes, upper-ocean mixing, MABL response to SST

80 fluctuations). They also have their unique aspects. Their specific processes that need special  
81 attention for the planning of TEPEX field observations are discussed below.

82 The cold tongue exists despite strong local solar warming because of intense upwelling of  
83 cold water and downward transport of heat by ocean mixing. This two-way communication of  
84 energy, momentum, and water properties between the sea surface and the thermocline couples  
85 the atmosphere to “ocean memory” that shapes ENSO. Air-sea and ocean processes determine  
86 the height, stability, and cloudiness of the local MABL, which determine how effectively the  
87 thermocline memory and MABL properties will be connected to the troposphere and thereby  
88 influence atmospheric general circulation. Locally, gradients of surface pressure - modulated by  
89 the SST gradient - competes with the tropospheric pressure gradient in driving the surface wind.  
90 The penetration of tropospheric momentum into the MABL and to the surface depends on MABL  
91 instability, which is mainly determined by SST. Local and remote factors modulating upwelling,  
92 and its role in regulating SST, air-sea fluxes, their spatial gradients, and their influence on stability,  
93 height, and cloudiness of the MABL are all largely unconstrained by *in-situ* or satellite  
94 observations, and therefore cannot be confidently modeled or simulated. TEPEX will provide the  
95 physical basis to underpin and advance the next generation of models and observing networks in  
96 the cold tongue.

97 The zonal movement of the warm pool at its eastern edge is forced by surface zonal winds  
98 and controlled by several factors (e.g., directly forced surface current, oceanic Kelvin waves,  
99 Ekman divergence, and the asymmetry in wind responses to the ocean). A special feature of the  
100 warm pool is the episodes of strong surface westerly wind events (WWEs) and strong rainfall  
101 events that are often closely related through, for example, the MJO. The counter-effects of  
102 enhancing ocean mixing by WWEs and weakening mixing by the barrier layer due to surface  
103 freshwater input by rain are perhaps among the most difficult to be accurately reproduced by  
104 numerical models. This difficulty is compounded by the multi-scale variability (the MJO, equatorial  
105 waves, extratropical intrusions) in the atmosphere, their interactions, and their ocean responses.  
106

## 107 **Hypotheses**

108 Decades of shipboard, moored, and satellite observations, as well as recent modeling  
109 results, have spurred new ideas for understanding and representing relationships between  
110 subsurface mixing, air-sea fluxes, surface forcing, and MABL response in the air-sea transition  
111 zone. They also prompted new hypotheses for the role of equatorial mixing and freshwater input  
112 in the Pacific coupled ocean-atmosphere system. The following hypothesized processes are  
113 proposed to guide TEPEX field observations:

### 114 115 *1) Upwelling and Mixing*

116 In the eastern Pacific cold tongue, persistent vertical shear between the surface westward  
117 South Equatorial Current and the eastward Equatorial Undercurrent in the thermocline creates an  
118 unstable regime below the surface mixed layer that is primed to mix vigorously. Surface-trapped  
119 currents in the afternoon warm layer set off a diurnal cycle of downward-propagating turbulence  
120 that reaches far below the ocean surface mixed layer into the upper thermocline. The resulting  
121 deep vertical mixing transports surface heat downward and cold water upward, cooling the  
122 surface and allowing the cold tongue to extend much further west than simple flow along the  
123 upward-sloping thermocline would imply. Thus, mixing drives efficient surface-thermocline  
124 communication even where the wind stirring does not reach the depth of the thermocline (Fig. 2).  
125 Mixing carries ocean memory to the atmosphere, enabling large-scale ocean-atmosphere  
126 coupling (Bjerknes feedbacks), and providing a path for carbon-rich water to reach the surface  
127 resulting in CO<sub>2</sub> outgassing, and sustaining the cold tongue’s productive fishery.

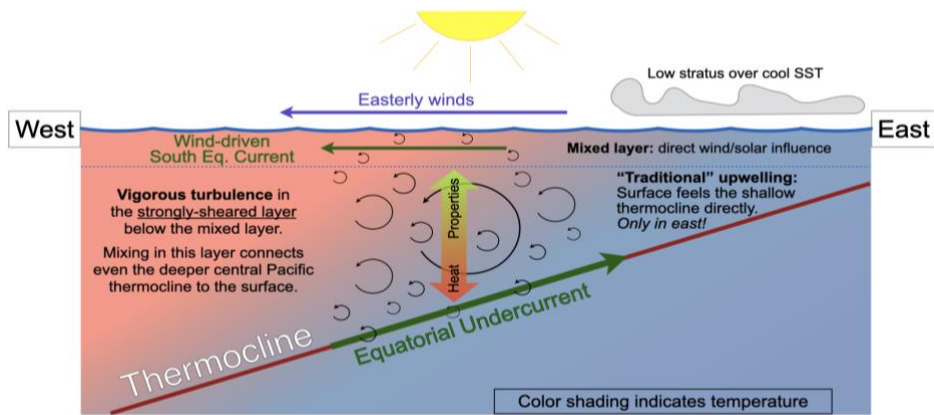


Figure 2 Schematic illustration of vertical-shear driven mixing in the cold tongue.

In contrast, in the region of the eastern edge of the warm pool, mixing is primarily driven by surface wind and inertial waves. Because of substantial surface freshwater due to rainfall, salt-stratified barrier layers can form, lasting until being eroded by subsequent strong surface wind forcing. Weakened or diminished mixing across the pycnocline increases the efficiency of wind-forced surface current responses.

## 2) Roles of surface freshwater input

In addition to forming salt-stratified barrier layers that inhibit entrainment of cold water from below and enhance surface warming, surface freshening from rainfall can also stabilize the upper ocean and generate density gradient-driven surface pressure jets along the equator that expand the warm pool eastward. The expanded warm pool can support further eastward propagation of subsequent atmospheric events, such as the MJO, and their associated westerly wind forcing, thus strengthening equatorial downwelling oceanic Kelvin waves that deepen the thermocline in the central and eastern Pacific. Such multiscale air-sea interactions, as illustrated in Fig. 3, are among the processes likely contributing to the onset of El Niño. However, this process is absent from the cold tongue region because of its weak precipitation.

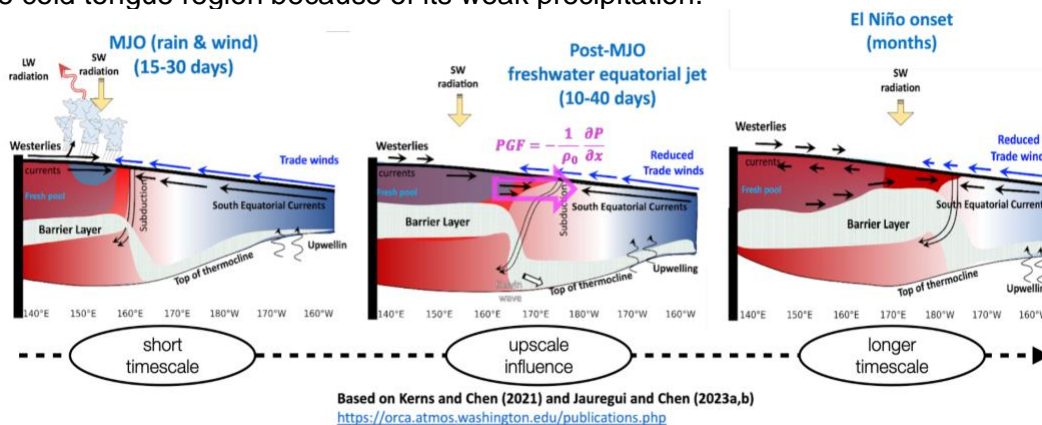


Figure 3 Schematic illustration of one of the hypothesized multiscale air-sea interactions of the warm pool. MJO precipitation and westerly winds induce eastward propagating oceanic Kelvin waves that deepen the thermocline and upper ocean barrier layer (left). During the post-MJO phase, surface warming and freshwater generate a large-scale zonal gradient in upper ocean pressure, which induces a strong easterly current (middle). Consequently, the warm pool is expanded eastward, and the trade wind is relaxed during the onset of El Niño (right).

155  
 156  
 157  
 158  
 159  
 160  
 161  
 162  
 163  
 164  
 165  
 166  
 167  
 168  
 169  
 170  
 171  
 172  
 173  
 174  
 175  
 176  
 177  
 178  
 179  
 180  
 181  
 182  
 183  
 184  
 185  
 186  
 187  
 188  
 189

3) *Air-sea coupling*

Air-sea coupling of the warm pool is empowered especially by its episodic, very strong precipitation events. Variations in surface wind, rainfall, radiation, and other quantities crucial to the atmospheric forcing of the ocean are often closely related to atmospheric deep convection. The oceanic feedback to the atmosphere through SST anomalies and air-sea fluxes is received by the MABL and manifested by the responses in atmospheric convection. In contrast, non-precipitating processes dominate air-sea coupling of the cold tongue. Variations in upwelling and mixing produce SST anomalies gradients on subseasonal to interannual scales. Through air-sea fluxes, atmospheric responses in near-surface moisture, temperature, pressure, boundary layer stability and height, and cloudiness determine the strength of turbulent mixing in the MABL. The consequential surface pressure gradient, tropospheric pressure gradient (measured at the top of the MABL), and cloud-enhanced vertical mixing of tropospheric momentum into the AMBL compete to drive surface winds, closing the interaction loop between the thermocline, air-sea transition zone, and the troposphere.

**Guiding Questions for TEPEX**

1. What controls the strength and extent of the thermocline variability and upper-ocean shear/stratification? What are the relative roles of upwelling, mixing, and surface forcing in this?
2. What processes control the emergence, strength, persistence, and pattern of SST anomalies? What are the relative roles of upwelling, mixing, thermocline anomalies, upper-ocean stratification, and surface forcing in this?
3. To what extent - and by what mechanisms - do changes in SST and air-sea fluxes, especially across the sharp front of the cold tongue and the eastern edge of the warm pool, co-evolve with the other properties in the air-sea transition zone? To what extent does the variability of the air-sea transition lead to the fluctuations in atmospheric convection and cloudiness and in the thermocline?
4. What are the roles of oceanic and atmospheric propagating phenomena (e.g., TIW, MJO, equatorial waves, inertial oscillations) in equatorial air-sea coupling? Do these processes connect the warm pool east edge with the cold tongue?
5. To what extent can the identified processes in the air-sea transition zone key to equatorial air-sea coupling be inferred or diagnosed from the sustained TPOS and routinely observed for accurate ENSO prediction?

**Field Observation Strategy**

<b>TEPEX-Central: eastern edge of warm pool</b> (across the 28.5°C isotherm, typically 160°E - 160°W, 5°S – 5°N)	<b>TEPEX-East: cold tongue</b> (140°W, 2°S – 5°N)
Boreal Spring (March – May) 2026	One year starting from boreal fall 2026 (three ship cruises, six-month apart)
Ship observations of the air-sea transition zone (including ocean mixing, convective clouds, and precipitation), the thermocline, and the tropospheric profiles	
UxS (underwater, surface, and aerial drones) observations of the air-sea transition zone	
Aircraft observations of the air-sea transition zone	enhanced moorings for ocean mixing
Land-based observations of the MABL and atmospheric convection	
Argo floats, surface drifters, TAO array	

Satellite observations of SST, SSS, wind, rainfall, cloudiness, water vapor, etc.

Real-time forecasts and state-estimation will guide the field operations and provide early interpretation and quantification of the value and impact of the field observations.

190

191

**Outcomes:**

192 a) **Advance fundamental understanding.** TEPEX will provide comprehensive, co-located in  
193 situ observations that will, through testing the proposed hypotheses, bring improved process-  
194 level understanding of the cold tongue and warm pool in the context of the air-sea transition  
195 zone across dynamical regimes; It will also pursue broaden scientific scopes into oceanic  
196 biogeochemistry and atmospheric chemistry;

197 b) **Guide model development.** TEPEX observations will provide the base for iterative  
198 implementation and evaluation of prediction model improvement: parameterizations of ocean  
199 vertical mixing and explicit coupling of freshwater and salinity, and for constraining large eddy  
200 simulations (LES) to produce observational proxies to assist model development;

201 c) **Guide observing system design.** TEPEX observations will help determine the variables and  
202 their spatial and temporal scales that must be sampled by the TPOS observing systems and  
203 satellite observations to produce adequate data for diagnostics of the air-sea transition zone  
204 over the cold tongue and warm pool and for accurate ENSO prediction.

205

206

**Impacts:**

207 The international TPOS 2020 observing system redesign identified the processes targeted  
208 by TEPEX as key to realizing the potential global prediction skill depending on intensely coupled  
209 tropical Pacific air-sea interaction. Observational, technical, theoretical, and modeling advances  
210 of recent years have prepared us to accomplish the goals laid out above. The insight gained  
211 through TEPEX co-located measurements of the air-sea transition zone and its interaction with  
212 the thermocline and troposphere will lead to more complete and more effective ocean-atmosphere  
213 monitoring of the region by the sustained observing system. The advanced observing system,  
214 understanding, and models in combination will lead to more accurate prediction of the Pacific  
215 ocean-atmosphere coupled system of a wide range of timescales. This advanced observation-  
216 understanding-prediction enterprise would have far-reaching societal benefits. It is the foundation  
217 for mitigation of the impact of extreme environmental events modulated by ENSO and the MJO  
218 (floods, droughts, wildfire, heat waves, marine heat waves, tornados, tropical cyclones, coastal  
219 inundation, etc.), better management of natural resources (fisheries, coral reefs, etc.), design of  
220 sound and sustainable climate solutions (clean energy, reforestation, aquaculture, CO<sub>2</sub> removal,  
221 etc.), and global/regional equality for vulnerable populations.

222 TEPEX will serve as the first field project of the WCRP Global Precipitation EXperiment  
223 (GPEX) and Precipitation Prediction Grand Challenge (PPGC), a WMO Ocean Decade Action.  
224 TEPEX will provide experience and lessons learned to pave the road for the long-term success  
225 of GPEX and PPGC.

226 TEPEX will promote diversity, inclusion, and equality by actively recruiting and supporting  
227 participants who are early-career and from underrepresented groups in geoscience to help build  
228 a more representative future workforce. TEPEX is envisioned as an international program that  
229 promotes collaborations and coordination among the countries/regions of the Pacific Rim to  
230 advance their common interests in the ability to predict the Pacific ocean and atmosphere to  
231 benefit society.